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Physical Properties of Rocks

6110 Elasticity, fracture, and flow

SURFACE CRACK GROWTH IN GEOLGICAL MATERIALS

B. F. Atkinson (Geology Department, Imperial College, London, U.K.)

A review is presented of the experimental data on subcritical crack growth in geological materials. The main parameters describing subcritical crack growth are the stress intensity factor (K_c), the subcritical crack velocity (V_c), and the stress intensity factor/crack velocity (K_c/V_c) relationship between K_c and V_c . The K_c/V_c data are presented in terms of an equation, which shows crack velocity depends on stress intensity factor rather than stress because this is common practice in experimental studies. The data are presented as tables and in synoptic diagrams. For all materials, the value of n increases as the environment becomes more aqueous and with increase in the microstructural complexity of the material. Values of n as low as 0.5 have been found for granite crazing in acidic aqueous environments, and as high as 170 for travertine crazing in water-saturated calcareous rocks.Insufficient experimental data are available to predict subcritical crack growth behavior at depth in the earth's crust without major extrapolation of the data from surface experiments. Therefore, the probable influence on subcritical crack growth of key parameters in the crustal environment. These include stress intensity factor, temperature, pressure, particle size, and presence of chemical agent, microstructure, and residual stresses. To predict crack growth in the subsurface, it is necessary to know the third derivative of the likely magnitude of the subcritical crack growth limit. For across-cortical crazing tests, values of n as high as 1000 and as low as 0.2 of the critical stress intensity factor are inferred from theoretical calculations. Further problems arise with regard to the extrapolation of experimental data to depths of 10 km. This includes the choice of a suitable equation to describe the crack growth behavior. A brief discussion of the double torsion testing method is provided in which the importance of the choice of experimental results because it is often the sole method used to study subcritical cracking in rocks, such as caving, crazing, fracture mechanics, extrapolation to natural conditions, etc.

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of the mechanical data, in combination with microstructural observations, are incompatible with existing models of subcritical crack growth behavior. Instead the strength/grain size relations are interpreted as resulting from variations in structurally incorporated water, which in turn are controlled by different mechanisms for the grain boundaries. Finally, contains areas of crenulations, shear bands, activation energy Q and stress exponent n for creep do not depend on grain size, so that the difference in strength must be incorporated in the stress intensity factor. However, both show a continuous decrease with increasing amounts of available water, from $Q = 300 \text{ kJ/mole}$ and $n = 4$ for water-soaked samples, to $Q = 130 \text{ kJ/mole}$ and $n = 4$ for dry samples.

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out of the depression is episodic in nature and has been inactive the last million years along the western boundary. To the east, there is clear evidence of uplift and lifting of sedimentary layers. A recent report has suggested that the current is responsible for the creation and cancellation of median valley relief. The transition between the upland and the trough is gradual and may take place within 1 to 2 km. The width of the transform fault just west of the depression is less than a kilometer. Hilltopping areas are located and displayed. Hilltopping is the diversion of surface runoff into hilltops, creating small lakes. Microcrusts are located by 3 or even 6 m above the ground. Aeolian activity occurs along the eastern wall of the median valley. In the eastern side, in the eastern part of the depression and in the coastal mountains. Very little activity is associated with the western edge of the transform depression and the trace of the transform fault.

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6170 Structure of the lithosphere

GEODESTIC INTERCITY OF THE TIGRIS-EUFRATES RIVER (SOUTH PACIFIC)

A. Caronova and K. Domini (Groupe de Recherches de Géodésie Spatiale, CNRS, 31025 Toulouse Cedex, France)

A comparison for the backscattered intensity of seismic waves, single scattered from a region containing fluctuations in the surface velocity has been derived for high frequencies by comparing the correlation of the slowest fluctuations in a Taylor series about zero lag. The resulting expression indicates that the backscattered intensity is independent of frequency and directly proportional to the first derivative of the autocorrelation at zero lag; the next term is proportional to the reciprocal of the square of the lag and directly proportional to the third derivative of the autocorrelation at zero lag. Contributions from terms of the Taylor series involving even numbered derivatives of the autocorrelation are negligible. Contributions from odd numbered derivatives of the autocorrelation are smooth function only at the smallest wavelengths and are zero at zero lag. This result demonstrates that backscattering at high frequencies can occur from fluctuations in velocity whose derivatives are opposed to fluctuations in the surface velocity.

The backscattered intensity is independent of frequency, the contribution of backscattering to the Attenuation factor is proportional to frequency.

Such behavior may have been observed in the ocean from the northeast (18°N - 23°N) as has been estimated. Together with dating data available for two seamounts of the chain, this result is available for the propagation of seismic waves compared to the visco-elastic model. For the flanking of the oceanic lithosphere in response to loading.

J. Geophys. Res., B, Paper 480149

2

6199 General

CONTINENTAL MARGINS, AND THE EXTENT AND NATURE OF THE CONTINENTAL SHELF

D. J. Thompson (Geography Department, Trent University, Peterborough, Ontario, Canada K9J 7B8)

A new map of the extent of continental crust shows that the continents are more extensive and thicker than previously thought. The total area is $10.4 \times 10^9 \text{ km}^2$, or 41% of the Earth's surface, and there are 14 of them. In order of decreasing size, they are: Russia, North America, Africa, South America, Antarctica, Australia, Asia, Mexico, New Zealand, Central America, Russia, Seychelles, Australia and Jan Mayen. Continental ocean contact is based on the basis of a literature review and empirical data. Isopleths are drawn along plate boundaries. The average thickness of continental crust is 36 km, small continents are over twice the large ones, and heights are significantly greater than thickness.

The thickness of the continental shelf is also significant.

There are several sources of error in the altimeter measurement system; some of these

The Potential of Satellite-Based Radar Altimeters

C. N. K. Mooers,¹ D. E. Barrick,² R. E. Cheney,³ D. B. Lame,⁴ and J. G. Marsh⁵

Introduction

Seasat is the only dedicated, multisensor oceanographic satellite to have ever flown. Although the mission lasted for only 3 months in 1978, it returned an enormous harvest of unique and valuable data that are still in the early stages of analysis. Its sensor complement included a scatterometer for surface winds, a passive microwave radiometer for sea surface thermal measurements, a synthetic aperture radar system for ocean surface imaging, and the radar altimeter. The altimeter provides information on significant wave height, surface wind speed, and above all, the topography of the sea surface. This topography is of prime interest to ocean circulation studies because its horizontal gradient, after removal of the geoidal signal, is proportional to the surface geostrophic current, the principal circulation component.

Satellite altimetry was introduced as a brief experiment on Skylab in 1973, followed by the 5.5 year GEOS-3 limited coverage mission which was launched in 1976. Seasat provided a superior altimeter system with decimeter accuracy and data recording capability which provided global coverage.

Looking to the future, the Navy plans to launch the lidar satellite GEOSAT in 1984 for geodetic applications. This mission has oceanography as a secondary objective.

The Navy is also planning NROSS (Navy Remote Ocean Satellite System), a Seasat type of mission, for 1988. The National Aeronautics and Space Administration (NASA) is preparing to launch TOPEX (Topography of the Ocean Experiment) in the same time frame with an orbit, a tracking system, and a highly accurate altimeter optimized for the study of ocean circulation. The European Space Agency is including an altimeter on the ERS-1 (European Remote Sensing) satellite in 1988, and the French Space Agency is considering the addition of an altimeter (POSEIDON) to their SPOT satellite in the late 1980s.

The Seasat Altimeter Data Seminar, held November 2-4, 1982, at the Jet Propulsion Laboratory (JPL), in Pasadena, Calif., brought together sensor and system engineers, geodesists, physical oceanographers, geophysicists, data managers, and others to consider the practical and scientific potential of satellite-based radar altimeters, especially those comparable to or better than Seasat. With much novel data in hand, and with the prospect of a major increase in activity in a few years, this seminar was held at a propitious moment in the history of ocean science.

The goals of the seminar were to review the information content of satellite altimetry for ocean science, to produce synergistic interactions in a unique aggregation of engineers and scientists, and to characterize the capabilities of satellite altimeters of the past and future. Ancillary objectives included (1) improving and broadening the use of satellite altimetry in ocean science for the betterment of the TOPEX program and (2) defining the needs of the scientific community for improved data base management for satellite altimetry through PODS (Pilot Ocean Data System). The purpose of this article is to provide a summary of these topics to the broadest possible scientific community.

Orbit Computation and Mean Sea Surfaces

There are four components to the determination of the satellite orbital path: (1) the differential equations which describe the motion of the satellites, (2) the solution of these differential equations, (3) the measurement of the satellite translational motion, and (4) the solution technique which enables estimation of the orbit, geodetic, geophysical, and oceanographic parameters from the first three components.

Tracking data are required to determine the orbit. Weighted least squares are used to solve for the position and velocity which best fit the tracking data. The differential equations of translational motion require accurate models of all forces acting on the satellite, including those of gravitational and nongravitational origin. Ground-based measurements also require modeling of the orientation of the earth in space with respect to the moon, sun, and planets and the orientation of the earth with respect to the angular velocity vector. The location of the tracking instrument on the earth is also required.

Precision orbit determination has been primarily limited by incomplete global tracking, errors in the model constants, or in the models of the physical forces acting on the satellite, most notably the earth's gravity model. The radial accuracy of the orbits contained in the Seasat geophysical data record (GDR) was approximately 1.5 m. A significant improvement in the earth's gravity model has since been achieved by incorporation of the Seasat altimeter data in the model adjustment process.

The altimeter data are especially useful since they are accurate and very dense coverage over the oceans is available. The improved model, called PGS-S4, has yielded radial orbit accuracy of better than 1 m when used with laser and Unified S-Band data in orbital lengths of several days.

Recent computations using the PGS-S4

Seasat data the altitude error was reduced to 6 cm.

Precision orbital computations and the altimeter data have been combined to produce global maps of the mean sea surface. The GEOS-3 and Seasat altimeter data have been used to compute a global representation of the mean sea surface on a $1^\circ \times 1^\circ$ grid. The Seasat ephemerides were based upon 6-day arcs of laser and unified S-band tracking data computed with the PGS-S4 gravity model. The GEOS-3 altimeter data were adjusted into a common reference system using a combination of laser reference orbits and crossing-arc adjustment techniques. In addition to the global solution, regional solutions have been computed in the North Atlantic, North Pacific, Bering Sea, and the Gulf of Mexico (Figure 1). These regional solutions have been gridded with horizontal resolutions as fine as 15 km; their precision is of the order of 10 cm.

Wind, Waves, and Tides

From the amplitude of the altimeter's mean return power waveform, wind speed can be estimated. Significant wave height can be estimated from the stretching of the altimeter return. Furthermore, minimum swell height can be determined by assuming that the measured wave height consists of energy due to

Article (cont. on p. 86)

MEAN SEA SURFACE
BASED UPON SEASAT AND GEOS-3
ALTIMETER DATA
1975-1978

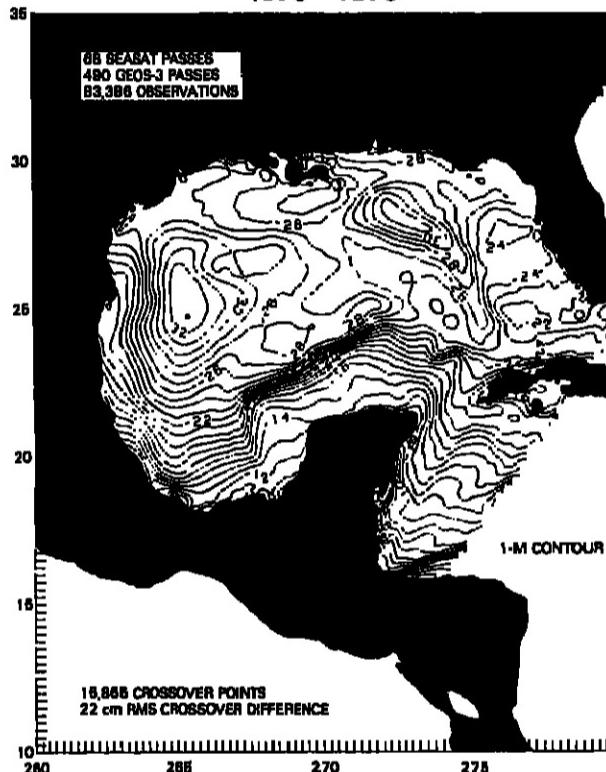


Fig. 1. Mean sea surface for the Gulf of Mexico based upon GEOS-3 and Seasat altimeter data of July-October 1978. The adjusted along-track data have been gridded on a $0.25^\circ \times 0.25^\circ$ grid [Marsh et al., 1984].

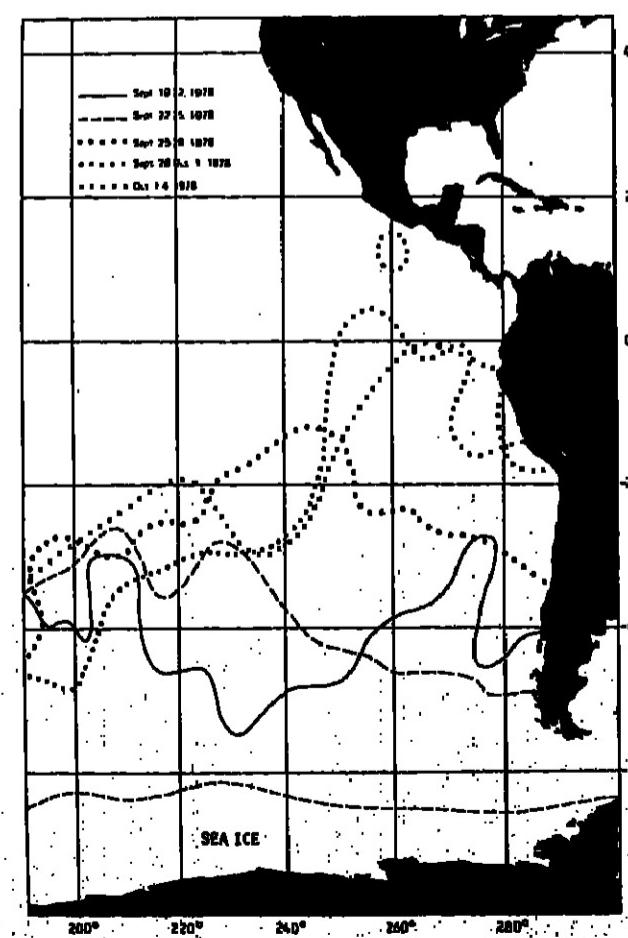


Fig. 2. Northward propagation of swell fields generated in the Antarctic as observed by the Seasat altimeter. Isopleths indicate the northernmost extent of swell greater than or equal to 5 m during the given 8-day periods [Mognard, 1984] (reprinted with permission, Crane, Russak & Co., New York).

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Lagrangian Studies of Deep Ocean Currents

T. Rossby

It is now more than 10 years since the first experiment using subsurface drifters (so-called Sofar floats) took place in the Mid-Ocean Dynamics Experiment (Mode). Since then they have been applied in a series of exploratory studies culminating in the Poly-mode Local Dynamics Experiment (LDE), and more recently along 55°W in a study of the North Atlantic subtropical gyre in a region well removed from the dynamics of western boundary regions. These neutrally buoyant instruments, which can be ballasted to drift with the water at pressures up to 2000 dbars (meters), are tracked acoustically over long periods of time (months to years). Conceptually the floats may be thought of as large molecules, fluid parcels whose pathways and speeds are explicitly known. The structure of their trajectories often yields surprisingly detailed information on the horizontal structure of the velocity field. When used in clusters they can tell us much about the mean field and the dispersive properties of the region. This article provides a brief retrospective of what we have learned in the 10 years since their first application in Mode. We begin with a brief description of the Sofar float technology.

Tracking Sofar floats over great horizontal distances is possible thanks to a remarkable acoustic property of the ocean known as the deep ocean sound channel or Sofar (sound fixing and ranging) channel. This acoustic wave guide, well known to acousticians and submariners since World War II, owes its existence to the happy fact that the speed of sound is a strong and comparable function of oceanic pressures and temperatures. In the upper ocean the speed of sound decreases rapidly with depth due to the thermal stratification; in the deep, nearly isothermal waters the speed of sound increases with pressure. The minimum speed of sound, about 1000–1300 m deep in the subtropical oceans is the axis of a permanent acoustic waveguide such that under quiet listening conditions one can hear a 1-W sound source at 250 Hz at distances greater than 1000 km.

The first suggestion to use the Sofar channel to track neutrally buoyant drifters was made by H. Stommel in 1949 in a paper on horizontal diffusion. In 1966 M. J. Tucker and D. C. Webb conducted an encouraging test of long-range transmission using a lightweight piezoelectric transducer, and in October 1969 a neutrally buoyant float was tracked for 4 months. But it was another 8 years before float could be put to use systematically.

Today's variety of float consists of an aluminum flotation tube, 0.3 m in diameter and 5.5 m in length, which also provides the housing for the battery pack and electronics. The transducer, a thin-walled tube 1.8 m long and of the same diameter, is open at one end and has a piezoelectric bender plate at the other. It is mounted end-to-end to the flotation tube.

Acoustic signals are transmitted every 8 hours; each signal consists of an 80-Hz FM pulse (1.5 Hz linear chirp) at 250 Hz. Besides giving a better signal-to-noise ratio than the previously used amplitude-modulated system, the phase modulation allows the use of simple digital correlators for signal detection and time of arrival determination. The radiated power levels have increased approximately from 3 W to 8 W, permitting tracking ranges out to 2500 km depending on the float's depth in the sound channel, and ambient noise conditions at the receiver site.

The floats are equipped with an active ballasting system to maintain a prescribed depth,

and telemetry of pressure and temperature. They are powered to last in excess of 2 years [Webb, 1977]. The early development of the Sofar float program was greatly simplified by the existence of landbased hydrophones on Bermuda, the Bahamas, and Puerto Rico. Their availability reduced the risk and cost of the program by permitting us to concentrate on the major technological uncertainty: the float itself.

Tracking the floats is conceptually very simple: Given knowledge of the speed of sound in the ocean and the time of arrival of signals at two or more receivers, one can determine from the intercept of circles (known travel times) or hyperbole (travel time differences) the position of a float to within a few kilometers. With the recent development of autonomous listening stations (ALS), which can be moored in the sound channel for a year at a time, Sofar float studies are no longer restricted to areas within range of land-based hydrophones. This has added great flexibility to their use.

There is no question that the most powerful attribute of drifters is the horizontal information that is so effortlessly provided—effortlessly in the sense that even a single instrument can suffice to lay bare the circular structure of a Gulf Stream Ring, show the path of the Gulf Stream as it is swept downstream, or reveal the constraints imposed on fluid motion by variable bathymetry. As they drift they in fact articulate specific pathways and rates of displacement of fluid parcels; information that cannot be obtained solely from the observed distribution of different water masses.

In what way would a tracer (or a potential anthropogenic pollutant) disperse, and how rapidly? The Sofar floats provide us with a natural tool to examine these kinds of questions. With an ensemble of trajectories one can start to construct statistical statements about mean flow and rates of dispersion and juxtapose these with classical water mass analyses. Let us attempt a simple illustration.

Between 1976 and 1980 we obtained nearly two dozen float trajectories at 700 m lasting 6 months or longer. They were set at various latitudes, mostly in the vicinity of 70°W. If one examines their position as a function of time one finds that the floats set north of 28°–29°N disperse to the west and north, become entrained into the Gulf Stream and are rapidly advected to the east. The ensemble of floats to the south of 28°N show evidence of a cyclonic circulation to the south and east.

In Figure 1 we show a sketch of the trajectories of floats at 700 m after they have been subjectively smoothed to remove mesoscale motions. It suggests that a tracer that is injected into the Gulf Stream recirculation system will be trapped and repeatedly recycled; it is not likely to be flushed to the south. We can compare this with the distribution of tritium (^{3}H) along 55°W, Figure 2. Note that at 700 m the ^{3}H does not penetrate south of 25°–30°N, and in the deep waters it is restricted to latitudes north of 50°N.

A simple 2-dimensional interpretation would suggest that the ^{3}H has diffused this far since it was injected into the oceans in the mid-60's. The circulation pattern inferred from the floats, on the other hand, indicate that the waters north of 28°N are trapped to the north thereof and subject to rapid recirculation, whereas waters to the south of 28°N are associated with a circulation on a much larger spatial scale and, hence, a longer time scale.

This comparison is, of course, incomplete and would be more effective if we could look at the source or inflow conditions to the east of the section. This has not been done. The point here is that even 2 dozen trajectories can provide valuable path and dispersion information. The idea of using floats as an integral part of modern water mass analysis is, however, still in its infancy.

The dispersion of the floats can be used to estimate eddy diffusivities. On time scales of the order of months, dispersion of floats is dominated by mesoscale eddy processes, especially in regions of high eddy kinetic energy. Moreover, the diffusivity is apparently linearly related to eddy kinetic energy over a wide range corresponding to an integral time scale of about 8 days. The physical basis for this relationship is unclear, but it certainly provides a simple means of parameterizing mesoscale eddy mixing in numerical studies. Dispersion tends to be isotropic when eddy kinetic energy levels are high, whereas in qui-

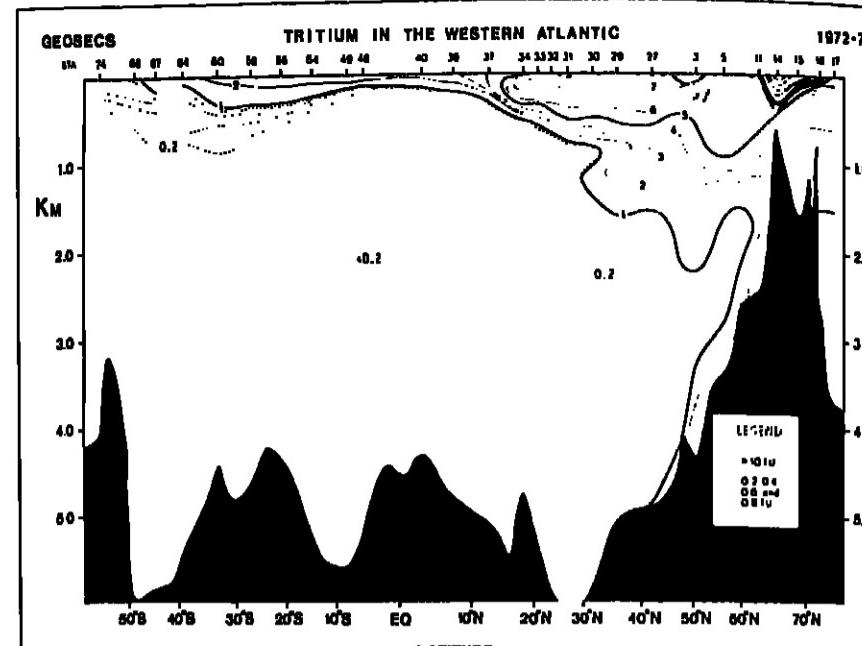


Fig. 2. Distribution of ^{3}H along a N-S section in the Western North Atlantic [Orlitz et al., 1977].

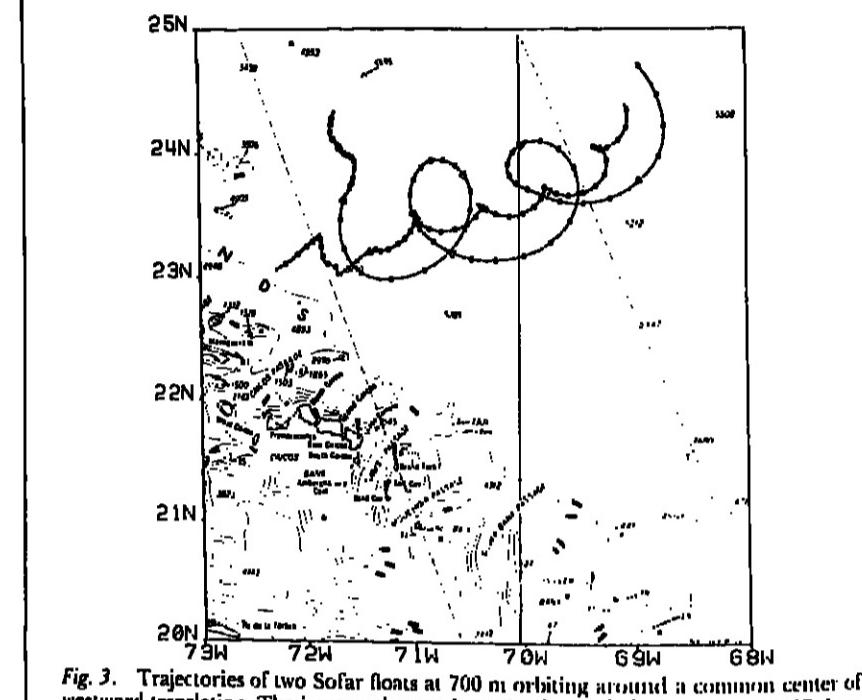


Fig. 3. Trajectories of two Sofar floats at 700 m orbiting around a common center of westward translation. The inner trajectory has a 10-day period, the outer one 17 days.

et regions such as the center of the subtropical gyre there is a clear tendency toward zonal dispersion.

On numerous occasions Sofar floats have exhibited astonishingly circular orbits with diameters ranging from a few kilometers to nearly 150 km. These trajectories reveal a class of oceanic motion that, apart from the ubiquitous Gulf Stream rings, was all but unknown 10 years ago. Spinning in either direction, these motions can be found in shallow as well as deep waters. And almost without exception their zonal motion is toward the west.

Figure 4 shows the trajectories of two Sofar floats (a third one is omitted for clarity) near 700 m depth spinning around a westward translating body of water until it appears to collide with the Bahama escarpment (or a boundary current along 10°) and the floats escape. Hydrographic observations at the time the floats were set revealed a thin lens of somewhat diluted Mediterranean water about 600 m thick and 120 km in diameter. Other observations of such lenses of Mediterranean water have since been made in the eastern Atlantic. Assuming a westward migration velocity of 3 cm s⁻¹ from its region of probable formation, the "needly," as it is sometimes referred to, must be at least 8 years old. With a mean period of revolution of 10 days, say, the lens must have made at least 100 revolutions since its genesis. Clearly the lenses are very stable, and as Figure 3 suggests, their demise may not be one of slow decay, but a sudden one due to changes in their environment, be it topographic "collisions" or their rupture by horizontal shear.

These lenses do not appear to have any atmospheric counterpart. What makes them particularly interesting is the suspicion that they may play an important role in the observed distribution of salt, oxygen, and other tracers in the ocean. Created in the east (it is not known how), they propagate zonally to the west and "deposit" the transported waters where they collapse. This suggests the possibility that observed distributions of water properties in some sense represent the probability density distribution of disintegration of these lenses and not solely a balance between a large-scale mean flow (to be determined) and eddy mixing, as is often assumed in diagnostic studies of ocean circulation.

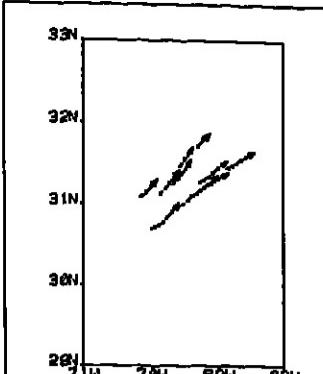
Conservation statements of the form $\frac{d}{dt}(*) = 0$ are intrinsically Lagrangian concepts where the property denoted by the asterisk remains invariant under translation of the fluid. The above mentioned "needly" is, of course, one example of fluid conservation. A corresponding dynamical test, namely the conservation of potential vorticity, has also been demonstrated.

Following a set of 10 Sofar floats for 2 months at 1300 m, Price and Rossby [1982] found that as the cluster moved north such that the local vertical component of planetary angular momentum increases, the cluster responds by turning in the opposite (or clockwise) direction so that its absolute angular momentum is conserved, and conversely so when they moved to the south (Figures 4a and 4b). For good numerical agreement it was found necessary to include vortex stretching caused by variable bathymetry.

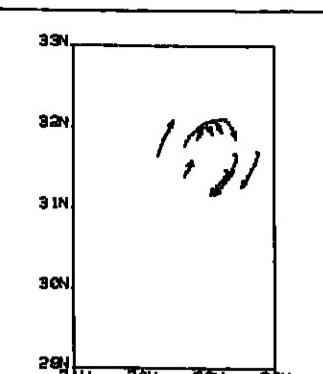
Horizontal arrays of Sofar floats have been used effectively to produce synoptic analyses of the velocity field. In both Mode and the Poly-mode LDE the systematic combination of velocity measurements at one level and hydrographic surveys have been used to analyze the dynamic state and evolution of the mesoscale eddy field. This methodology works well, but the rapid dispersion of floats limits to a few weeks the time during which accurate synoptic maps of the stream function can be constructed. The longevity of the Sofar floats is of no help here; thus, for future studies of this kind there may be a need for a simple, low-cost Sofar float of medium range of the kind used in the French Tourbillon experiment [Groupe Tourbillon, 1985]. For a more detailed discussion of the above ideas the reader is encouraged to consult chapters 4 and 5 in Robinson [1982].

Sofar floats are now in use in four experiments in the North Atlantic: The URI-WHOI Line and Gulf Stream Recirculation Experiments, Dispersion Studies in Very Energetic Eddy Fields (WHOI), Topogulf (a French program), and a nascent British study of dispersion in very deep waters. The technology is mature and reliable. The inability to track the floats in real time, however, has handicapped experimental plans to use floats interactively (for example in hydrographic surveys). This has stimulated a program called Relays (WHOI) to develop listening systems which are suspended down into the sound channel from drifting surface platforms. Tracking data can thus be relayed immediately to ARGOS (a satellite-based platform location and data collection system).

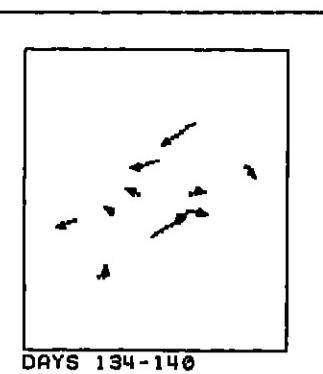
At present long-range float tracking is limited to oceans with permanent thermoclines where acoustic energy is trapped by refraction only. However, I believe it is possible to extend the technique to cold oceans with so-called half sound channels, i.e., where rays undergo surface reflection. For this to work,



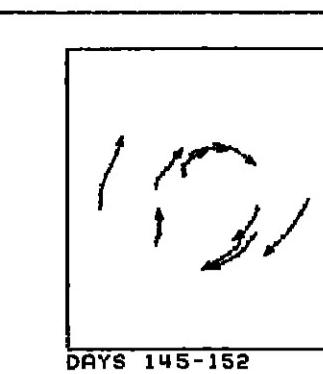
DAYS 134-140



DAYS 145-152



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DAYS 169-176

Fig. 4a. The geographical position of 10 floats at 1300 m in 12-day steps.

Fig. 4b. The same floats as in Figure 4a, but shown relative to the center of gravity of the cluster. Note the anticyclonic movement while the cluster is in the north, and the converse in the south [Rossby, 1982] (reprinted with permission, Springer-Verlag, New York).

observing small-scale processes without the problems associated with advection past stationary platforms.

Indeed, with the ready availability of sophisticated yet low-powered microprocessors, one can foresee the development of a variety of "intelligent" drifters designed to monitor the internal wave field, chemical changes due to isentropic mixing, or listen to and observe the local ecosystem. Days, weeks, or months later they can surface and report their findings; these can then be related to the large-scale processes within which they were embedded.

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Meeting Report

Quantifying Submarine Hydrothermal Fluxes

Many oceanographers believe that the chemical fluxes associated with deep sea hydrothermal processes are large and geochemically important, but quantifying these fluxes is proving difficult. Seawater-basalt exchange takes place in high-temperature hydrothermal systems at the very axis of seafloor spreading, as observed at the Galapagos Spreading Center [Cordts et al., 1979], and the East Pacific Rise at 13°N and 21°N [Murchie et al., 1982; RISE Project Group, 1980]. It also takes place in the lower temperature hydrothermal systems which are ubiquitous on the flanks of mid-ocean ridges, until sedimentation and void-filling in basalt, seal the hydrothermal systems at crustal ages between 10 and 80 m.y., e.g., Anderson et al., 1977.]

Chemical fluxes associated with crustal water/rock reactions have been estimated in two ways. One involves using the heat balance to estimate the rate at which hydrothermal solution exit the crust, and taking the product of this number and change in chemical concentrations to get chemical fluxes. The composition of hydrothermal solutions exiting at the ridge crest is well known; solutions are depleted in Mg²⁺ and SO₄²⁻, and enriched in Ca, alkalis, and heavy metals [e.g., Edmond et al., 1979; Mauz and Holland, 1978]. However, the heat flux is poorly known, with current estimates ranging between 0.2–5 $\times 10^{10}$ cal per year (0.8–21 $\times 10^{10}$ J per year) [Sleep et al., 1984; Edmond, 1981]. On the ridge flank, the opposite situation obtains. The convective heat flux is known to be about 5 $\times 10^{10}$ cal per year [Anderson et al., 1977; Williams and Von Herzen, 1974], but almost nothing is known of the composition of the reacted seawater.

The alternative method of constraining hydrothermal fluxes comes from the complementary approach of studying crustal chemistry. The composition of unaltered crustal rocks is known from detailed studies of few DSDP holes, and hydrothermal fluxes may be estimated from the crustal generation rate

and the crustal chemistry of representative sections [Hart and Staudigel, 1982; Thompson, 1984]. This approach is bearing fruit, but has a limited applicability because of the limited availability of samples taken within the oceanic crust.

To consider these problems, the University of Rhode Island's Graduate School of Oceanography recently hosted a 1-day symposium, under the auspices of the Norman Watkins Lecture Series, on the topic, "Quantifying Submarine Hydrothermal Fluxes: Evidence From Different Perspectives." The lectures of the seven speakers were discussed by members of the audience from URI and her sister institutions in New York and New England.

Norman Sleep (Stanford) and William Jenkins (WHOI) discussed constraints on the axial heat flux from thermal modeling and the oceanic He isotope distribution, respectively. Sleep estimated the maximum axial convective heat flux as 2×10^{10} cal per year (determined from the product of the total crustal generation rate and the latent sensible heat loss per gram of crust [Sleep et al., 1984]). He

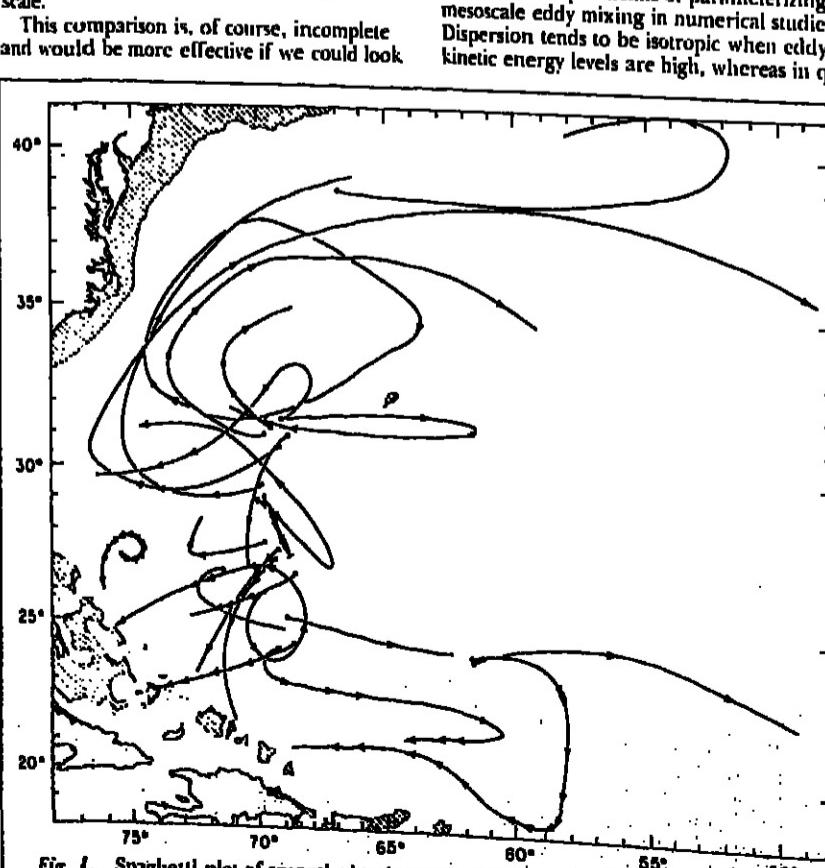


Fig. 5. Spaghetti plot of smoothed trajectories of Sofar floats at 700 m. Arrows are 100 m apart. Note the high velocity of floats caught in the Gulf Stream (from chapter 4, Rob-1983), reprinted with permission, Springer-Verlag, New York.

News

No Olivine in the Mantle?

Perhaps the most impressive factors in D.L. Anderson's analysis of new physical models of the earth are contributions from the numerous disciplines of modern geophysics, including 3-dimensional seismological observations, high-pressure experiments, highly precise isotope analyses, and studies of other solar system bodies (*Science*, 223, pp. 547-555, 1984).

The results? In short, there are the "ins" and the "outs." For example, the basalt-eclogite transition is back in fashion, whereas the notion of an olivine-rich deep mantle assemblage is no longer in fashion. This analogy is not to be construed as any return to old, pre-plate-tectonic concepts. Modern research, in the purest sense, is forcing a reexamination of some long-held assumptions.

Anderson would not only ask to have the concept of the basalt-eclogite transition revisited as a dominant crust-mantle parameter, but he provides insight to his suspicion of the validity of the olivine-spinel phase change and other olivine-related transitions as important boundaries in the transition zone. Thermal expansion and other thermally derived processes in the mantle (i.e., a hot, low-velocity zone) are suspect as well. So, therefore, the olivine-spinel transition is "out," as is the concept of a partial-melt, spherical shell, low-velocity zone from which bands could be derived. Instead, "the buoyancy differential that drives mantle convection is provided by partial melting and the basalt-eclogite phase change rather than thermal expansion" and "the large density changes associated with phase changes and melting in the basalt-eclogite system may drive convection and be responsible for the chemical stratification of the mantle and the long-term isolation of geochemical reservoirs." The result is that an olivine-rich mantle concept would not be compatible.

Anderson's synthesis of observational and experimental data, trace element analyses, and the approximations and model functions required to fill the gaps of knowledge is a courageous, and, of course, controversial, attempt toward taking an imaginative look at all approaches to deriving a meaningful earth model. It may be argued that the new interpretations of this model are no better than those existing, because many of the advances in seismology and experimental research on which the model is based are too new. Considerably more data are needed to justify many of Anderson's conclusions, and some large knowledge gaps will not be filled soon. Anderson cannot be faulted, however, for lack of imagination nor for creating a set of ideas presented in a scholarly way. Anderson's model should stimulate a strong response; the response may be in the form of obtaining the needed data.—PMB

Radio Telescope Center Selected

Socorro, N. Mex., will be home for the operations center for the Very Long Baseline Array (VLBA) network of radio telescopes. The National Radio Astronomy Observatory (NRAO) selected Socorro because of its proximity to the Very Large Array (VLA), an existing system of 27 radio telescopes and will allow combined operation of the VLA and the VLBA. In addition, two of the proposed VLBA antennas will be nearby. With the proposed array of 10 radio telescopes—from Puerto Rico to Hawaii—astronomers will be able to probe the universe with a resolution 1000 times greater than any existing radio or optical telescope and 100 times that of the future Hubble Space Telescope.

The VLBA will be operated by NRAO as a national facility. NRAO is operated by Associated Universities, Inc., a consortium of nine member universities under contract with the National Science Foundation. The proposed funding increase for VLBA for fiscal 1985 (up 82.1% to \$35.4 million) would go for construction of the array (EOS, February 14, 1984, p. 49).

Acid Rain Trends Summarized

In the northeastern United States, the acidity of precipitation has changed little in recent years, although the acidity is increasing in other regions. That's the latest word from a comprehensive review by the U.S. Geological Survey (USGS) of more than 200 published reports of acid rain research from the past 30 years. The report contributes to the controversy over whether increased sulfur emissions from Midwestern powerplants increase the acidity of precipitation in the Northeast.

"When the results of the many individual studies are combined, they show that acidification of precipitation in the Northeast-

which has the most damaging level of acidity on a regional basis, occurred primarily before the mid-1950's and has been largely stabilized since the mid-1960's," said John T. Turk, a research hydrologist at the USGS Denver office and author of the 18-page summary report.

Turk concluded that surface waters in lakes and streams in the Northeast follow a pattern of acidification similar to that of precipitation. The acidification of surface waters occurred before the mid-to-late 1960's; since then, some waters have not acidified further, and other streams show a slight recovery.

Trends in the acidity of precipitation in the southeastern and western parts of the country are far less certain. "In the southeastern United States, the available data show that precipitation is more acidic than would be expected for sites unaffected by manmade emissions," Turk said. "In addition, a comparison of recent precipitation data with the meager historical data suggests an increase in acidification of precipitation since the 1950's." Turk found, however, that most of the available data are ambiguous as to whether acidification of surface water has occurred in the southeast.

Copies of *An Evaluation of Trends in the Acidification of Surface Water in North America* (USGS Water Supply Paper 2249) are available for \$2.75 each from the Branch of Distribution, Text Products Sections, USGS, 604 S. Pickett St., Alexandria, VA 22304.

In Congress

Upcoming Hearings

The following hearings have been tentatively scheduled for the coming weeks by the Senate. Dates and times should be verified with the committee or subcommittee holding the hearing or markup; all offices on Capitol Hill may be reached by telephoning 202-224-3211.

March 7, Monday, March 14: Clean Air Act (P.L. 95-95) amendments (S.768) markup by the Senate Environment and Public Works Committee, Dirksen Building, Room SD-406, 10 A.M.

March 19: National Oceanic and Atmospheric Administration fiscal 1985 budget hearings by the Commerce, Justice, State, Judiciary, and Related Agencies Subcommittee of the Senate Appropriations Committee, Capitol, Room S-146, 2 P.M.—87R

Geophysical Events

Volcanic Events

Campi Flegrei (Italy): Uplift and seismicity in the caldera since mid 1982. Etna (Sicily): Incandescent tephra from central crater; seismicity.

Kilauea (Hawaii): 13th-15th major phases of East Rift Zone eruption include lava fountains to 300 m and temperatures to 114°C. Mt. St. Helens (Washington): Deformation and seismicity, then new lobe. Veniaminof (Alaska): lava fountains and flow continue.

Pavlof (Alaska): Plumes on satellite imagery harmonic tremor.

Piton de la Fournaise (Réunion Is.): Second phase of lava emission.

Sakurajima (Japan): 1988 explosions and ash falls tabulated.

Kusatsu-Shirane (Japan): 1988 activity summarized.

Rabaul (New Britain): Marked increase in unrest.

Manam (Bismarck Sea): Strombolian activity; explosion cloud to 3.5 km.

Langila (New Britain): Vulcanian explosions; ashfalls on coast.

Bagana (Solomon Islands): Two active lava flows.

Erebus (Antarctica): Seismicity normal; SO₂ flux measured.

Aerospheric Effects: El Chichón cloud persists; lidar data to north pole.

Campi Flegrei, S Italy (40.83°N, 14.14°E).

The following report is from Giuseppe Luongo, Roberto Scandone, and Franco Barberi: "Campi Flegrei (Phlegraean Fields) is a large caldera some 12–14 km across, located roughly 25 km W of Vesuvius and 15 km WSW of the city of Naples. The caldera formed after a huge eruption 35,000 years ago that produced 80 km³ of dense rock. Several other eruptions of decreasing intensity have occurred since then. In the past 10,000 years at least 22 different centers are recognizable. The last eruption occurred in 1858.

"Campi Flegrei has been the site of slow vertical movements since at least Roman times. A slow subsidence had occurred since the last eruption in 1858. An uplift that was observed in 1970 continued until 1972 without significant seismic activity. The inferred maximum uplift with respect to previous leveling was 170 cm. Slow oscillations of the

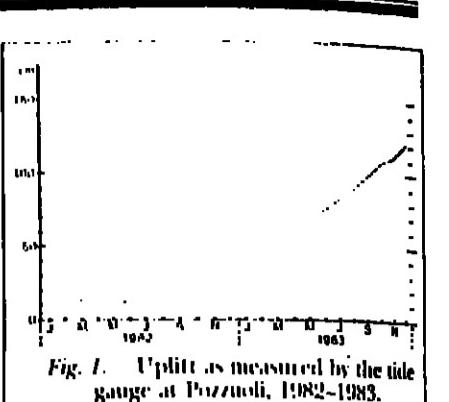


Fig. 1. Uplift as measured by the tide gauge at Pozzuoli, 1982–1983.

authorities in order to minimize the social consequences of evacuating people from their residences. The new settlement is relatively safe from a seismic point of view but is not safe from a maximum probable volcano event.

"In November 1982, moderate seismic activity was observed by the permanent seismic network, which has been operating since 1972. The level of activity was slightly above the microseismic activity in the area. In January 1983, public officials were notified of the anomalous trend of the phenomenon and of the possibility of an increasing seismic and volcanic hazard. In March a distinct increase in seismic activity was observed with the first magnitude 3 earthquake. Since then, ground uplift has continued with a velocity that reached 5 mm per day during October. After October, oscillations in the rate of uplift were observed, with a range between 1 and 4 mm per day. The seismic activity increased following a trend similar to that of the uplift velocity (Figure 2). A magnitude 4 earthquake occurred October 4, 1983, and its epicenter was in the Solfatara area. A close correlation seems to exist between the velocity of uplift and the seismic activity. The more energetic earthquakes seem to coincide with the higher rates of uplift (1–5 mm per day). The shallow character of the seismic activity does not give any evidence of a zone of anomalous propagation of S waves."

"Since April 1983, radon measurements have been made in water wells located in the

Correction

On p. 66 of the Feb. 21, 1984, issue of *EOS*, the paragraph in column 2 under *Other Awards* headed "Robert E. Horton Medal" should have appeared at the bottom of column 1, just below the section headed "Robert E. Horton Research Grant."

March 7, Monday, March 14: Clean Air Act (P.L. 95-95) amendments (S.768) markup by the Senate Environment and Public Works Committee, Dirksen Building, Room SD-406, 10 A.M.

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Books

Negev: Land, Water, and Life in a Desert Environment

Daniel Hillel, Praeger, New York, xx + 270 pp., 1982, \$32.95.

Reviewed by William Bach

In view of the continuing increased concern about the extreme fragility of deserts and desert margins, *Negev* provides a timely discussion of land-use practices compatible with the often conflicting goals of preservation and development. The success of agricultural and hydrological experiments in the Negev desert of Israel offers hope to the large percentage of the world's population that lives in an acceptably low quality of life in desert margins. Deserts are the one remaining type of open space that, with proper use, has the potential for alleviating the misery often associated with expanding populations.

This extremely well-written, entertaining book contains flashes of humor. It reads like a novel and is to a large extent autobiographical. The author has the uncommon talent to weave anecdotes into scientific facts and interpretation. The semitechnical style of writing, with minimal references, and the scrapbook nature of the photographs add readability and poignancy to the book. The first part (66 pages) is a perceptive description of the ecology of deserts that includes discussions of water, soil, vegetation, ecosystem, animals, and man's relation to the desert. The second part of the book relies heavily on the author's scientific work and personal experiences that are used to describe these elements within the context of the Negev, somewhat as a case study but, perhaps, more as a microcosm of this pioneering effort.

Another fascinating example of the relationship between archeology and hydrology, in addition to the author's comments on reconstruction and importance of the cisterns, is the explanation of the countless heaps, mounds, and strips of gravel found on many hillsides covering scores of square kilometers that are commonly arranged to form regular geometric patterns. Previous speculations were (1) that the gravel mounds supported grapevines rooted under them and that the heat emission from the dark gravel hastened the ripening of the grapes; (2) that the gravel mounds were "aerial wells" designed to condense dew during the night to irrigate the roots of the grapevines; and (3) that the mounds were built to increase the rate of erosion of the hillsides to hasten the deposition of soil in the bottom-land terraces.

Daniel Hillel, the internationally recognized soil physicist, was one of the original 12 founders of the pioneer settlement of Kibbutz Sdeh-Boker in 1951. Much of their early work was to study the methodology that early civilizations used to obtain water, to repeat their techniques, and to attempt to improve them. The author develops a strong archeological and historical theme about use of water from the time of the earliest people of the Negev up to the recent, sophisticated techniques of water management.

The main guides that he and his colleagues had for the Negev in the early days were the Bible and *The Wilderness of Zin* written by two British archeologists, C. L. Wooley and T. E. Lawrence, the latter who was to become the famed "Lawrence of Arabia." Nelson Glueck,

and monitoring of gas content and temperatures of fumaroles. Vertical ground deformation is measured by a repeated leveling of a permanent network and is also checked daily by a tide gauge in Pozzuoli harbor. The permanent seismic network operating in the area (Figure 3) is composed of 22 vertical seismometers, 15 of which are cable connected to a central point in Naples. A seismic explosion campaign has been planned in the Gulf of Pozzuoli to provide information on the deeper structure of the area. In cooperation with University of Wisconsin seismologists, a temporary network of 10 three-component stations with high dynamic range has been deployed in the area and will operate for some months.

Information contacts: Giuseppe Luongo and Roberto Scandone, Osservatorio Vesuviano, Largo S. Marcellino 10, 80138 Napoli, Italy; Franco Barberi, Dipartimento di Scienze della Terra, Via S. Maria 55, Pisa, Italy.

Rabaul Caldera, New Britain Island, Papua New Guinea (4.27°S, 132.20°E).

This report is from P. Lowenstein: "There was a marked increase in the amount of unrest in Rabaul Caldera during January, with a total of 8372 volcanic earthquakes recorded, an increase of 1255 over the December total (see last month's Bulletin)."

"A major seismic crisis took place on January 15 when 942 earthquakes occurred, including several strongly felt events. The maximum magnitude earthquake (ML 4.9) was accompanied by underground rumbling sounds. This crisis was accompanied by a maximum tilt change of 32.5 microradians at Rapindik tilt station and a lateral intrusion of about 0.3 × 100 m² of magma at a depth of 0.9–1.1 km. This resulted in a shift of the center of maximum uplift of about 1 km to the NW of its previous location, bringing it closer to Rapindik than to Tavurvur."

"The overall distribution of earthquakes in January was similar to that in December, with high concentrations on the NE (Greet Harbour) and W (Keravia Bay) sides of the harbor. Local concentrations of events also occurred along the Rapindik NE-SW fault line after the seismic crisis on January 15.

"Steady inflation of the Keravia Bay and Greet Harbour magma reservoirs continued throughout the month. The lateral intrusion of magma under Greet Harbour resulted in a maximum uplift of 6 cm in Greet Harbour and a vertical displacement of 3 cm along the Rapindik Fault."

"As a result of the increased activity in January, a warning was issued to the authorities to the effect that the eruption, which was previously thought to be only a possibility when the stage 2 volcanic alert was declared on October 28, was now much more likely to occur within the next few months."

Information contact: P. Lowenstein, Principal Government Volcanologist, Rabaul Volcano Observatory, P.O. Box 386, Rabaul, Papua New Guinea.

Earthquakes

Information contacts: National Earthquake Information Service, U.S. Geological Survey, Stop 967, Denver Federal Center, Box 25046, Denver, CO 80225.

Meteoritic Events

Fireballs: Austria (2); Austria-Czechoslovakia; N Central USA—S Central Canada; Florida, Pennsylvania—New Jersey, Pennsylvania, USA

This is a summary of *SEAN Bulletin*, 9(1), January 31, 1984, a publication of the Smithsonian Institution's Scientific Event Alert Network. The entire Rabaul article is included: the Campi Flegrei and earthquake reports are excerpts. The complete bulletin is available in the microfiche edition of *EOS* as a microfiche supplement or as a paper reprint. For the microfiche, order document E94-002 at \$2.50 (U.S.) from AGU Fulfillment, 2009 Florida Avenue, N.W., Washington, DC 20009. For the paper reprint, order *SEAN Bulletin* (giving volume and issue numbers and issue date) through AGU Separates at the above address; the price is \$3.50 for one copy of each issue number; for those who do not have a deposit account, \$2 for those who do; additional copies of each issue number are \$1. Subscriptions to *SEAN Bulletin* are available from AGU Fulfillment at the above address; the price is \$18 for 12 monthly issues mailed to a U.S. address; \$28 if mailed elsewhere, and must be prepaid.

Earthquakes

Date	Time, LST	Magnitude	Latitude	Longitude	Depth of Focus	Region
Jan. 1	0904	6.5 mb	33.40°N	137.32°E	37.1 km	S of Honshu, Japan
Jan. 8	1524	6.6 M	2.82°S	118.80°E	shallow	W Sulawesi, Indonesia

Now Available...

Journal of Atmospheric Chemistry

Editors:
P. J. CRUTZEN, Max-Planck-Institute for Chemistry, Mainz, F.R.G.
D. H. EHHLALT, Kernforschungsanlage Jülich, F.R.G.

The *Journal of Atmospheric Chemistry* is devoted to the study of the chemistry of the Earth's atmosphere with emphasis on the region below about 100 km. This field of research has grown vastly over the last decade, especially after the recognition that Man's activities can influence important processes of atmospheric chemistry on a global scale. Atmospheric chemistry is strongly interdisciplinary and embraces such sciences as chemistry, physics, meteorology, oceanography, soil science, biology and microbiology. For this reason, reports on research in atmospheric chemistry have been rather scattered in the literature. It is hoped that this new journal will fill a role in concentrating the flow of information. The Journal publishes original work on: observational, interpretative and modelling studies of composition and physico-chemical processes in the Earth's atmosphere, excluding air pollution problems of only local importance; the role of the atmosphere in biogeochemical cycles; chemical interaction of ocean, land surface, biosphere with the atmosphere; laboratory studies of the mechanics in homogeneous and heterogeneous transformation processes in the atmosphere; description of major advances in instrumentation developed for the measurement of atmospheric composition and chemical properties.

Subscription Information

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Article (cont. from p. 81)

swell plus wind waves in fully developed seas. Monthly global maps of wind, significant wave height, and minimum swell from Seasat show not only the expected zonal pattern due to the trade winds and other major wind systems but also wind and wave features on scales as small as 1000 km. Winds and sea states were highest in the Southern Ocean, and the local maxima migrated eastward from the Atlantic to the Indian Ocean and finally into the Pacific during the summer of 1978. Using successive 3-day maps, swell fields have been tracked from their initial formation in the Southern Ocean northward through the Pacific toward North America (Figure 2).

Amplitudes and phases of ocean tidal components can be recovered with satellite altimeter data. The amplitude and phase of the M₂ tide in the Indian Ocean were obtained from a 2-dimensional space-time least-squares harmonic analysis of the last month of the collinear Seasat data, Figure 3. There are four interacting amphidromic points surrounding a large area of maximum amplitude and stationary phase. This solution shows an enhancement of 10 to 20 cm in the maximum amplitude in the middle of the ocean compared with most models. Comparisons with the Schwiderski model indicate a shift northward for the amphidromic point near Australia and a southward shift for the one near Madagascar.

Ocean Circulation and Variability

Mesoscale eddies (scales of 50 to several hundred kilometers) occur in all oceans and are responsible for much of the horizontal mixing. The most intense eddies are associated with western boundary currents and other concentrated flows. Eddies may alter the amplitude of the sea surface by as much as 1 m. The meandering of intense currents, which generate many of the eddies, is considered part of the total eddy field.

Many of the altimetric techniques developed for observing sea height variability due to eddies are independent of orbit and geoid error. The method of collinear differences can be used for the last 25 days of the Seasat mission when the ground track was repeated within 2 km every 3 days. Mesoscale variability can be observed in those repeated profiles since the geoid is constant in time. Meandering currents and migrating eddies appear as wave-like signals propagating through the altimeter profiles. A global map of mesoscale variability compiled from all the Seasat collinear data (cover) shows the largest variability associated with five major current systems: the Gulf Stream, Kuroshio, Agulhas, Antarctic Circumpolar, and the Falkland/Brazil confluence. As expected, there is a marked contrast between high energy in the western parts of ocean basins and low energy in the east. Several areas such as the Eastern Pacific and South Atlantic are remarkably quiet with rms variability of only 1–2 cm. Because of the existence of these vast, low-energy regions, the North Equatorial Current systems in both the Pacific and Atlantic appear zonal variability maxima. The relationship of this variability to bottom topography in the Southern Ocean can be interpreted in a manner consistent with theoretical concepts. For example, the generation of anticyclonic eddies is suggested downstream of the Macquarie Ridge.

Using this collinear data set, the wavenumber spectra of mesoscale variability were found to be a function of energy level. In days of Seasat altimeter data from which a Fourier series representation of the radial orbit error was removed through an analysis of crossover differences (where ascending and descending ground tracks intersect), together with the along-track altimeter data. A third surface was determined from a spherical harmonic analysis of the differences between the GEM-9 geoid and a 3-month Seasat altimeter surface of the Pacific. All three of these surfaces show similar gyre-scale features, remarkable result considering the magnitude of the present altimetric and geoidal errors.

Ocean Models

A quasi-geostrophic, statistical-dynamical model has been used to simulate a satellite altimetric mission using the POLYMODE data set. The net improvement of ocean predictions achieved by 4-dimensional data assimilation of altimetric data, in addition to in situ observations, through optimal estimation theory, was demonstrated for hypothetical satellite tracks. Satellite altimetry can be further used to develop the correlation functions needed on a global basis to specify the error model for that method.

Eddy-resolving general ocean circulation models demonstrate the spin-up of an oceanic gyre by steady wind forcing and the generation of eddies by various hydrodynamic instabilities. The models reach a dynamical equilibrium between the mean and fluctuating flows, whose statistics show reasonable agreement with observations. These results suggest productive simulations could be performed in advance of an altimetric mission to determine whether eddy generation processes could be tracked by satellite altimetry. Also, statistics determined from altimetry could be used to validate the models.

A two-layered, semi-enclosed basin (Gulf of Mexico) model simulated that the eddy-shedding process could be studied by satellite altimetry. Due to two eddies are shed per year, with horizontal scales of 50 to 350 km and dynamic height amplitudes of 75 cm. Seasat collinear passes have produced altimetric signatures consistent with model predictions. An equatorial reduced gravity wave model driven by slowly varying wind forcing has demonstrated the significant, wind-driven circulation events which propagate along the equatorial waveguide, e.g., those which lead to El Niño over the course of a few months. Satellite altimetry may contribute to studying this problem, and this model seems promising for sensitivity testing of the equatorial oceanic response to the quality of the atmospheric forcing determined from conventional wind data versus "unconventional" satellite wind data sources. A two-layered, global scale, dynamic general ocean circulation model is being developed by the Naval Ocean Research and Development Activity for operational use at the Fleet Numerical Oceanography Center. A preliminary, one-layered version yields an

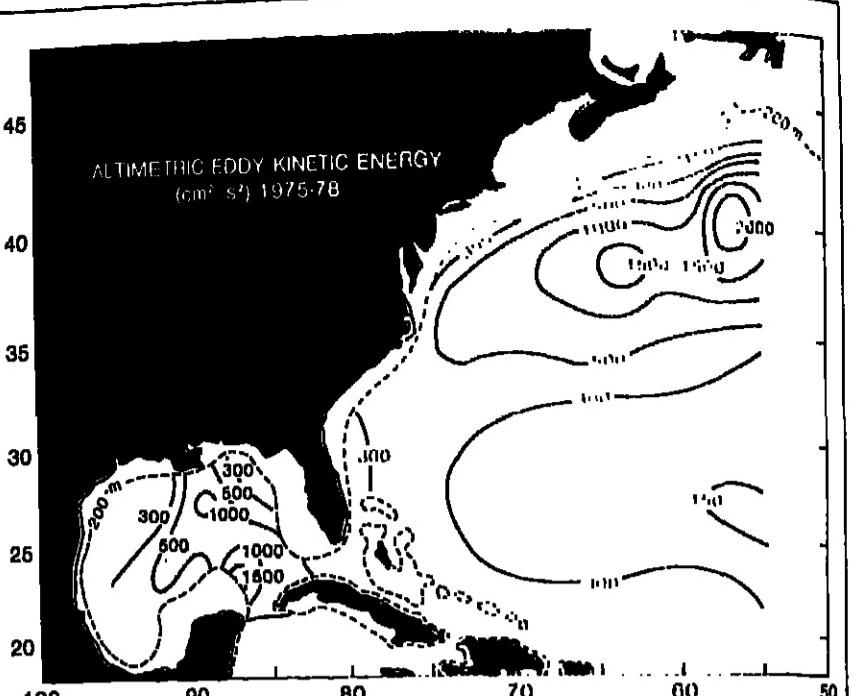


Fig. 4. Eddy kinetic energy computed from GEOS-3 altimeter over a period of 3.5 years by the collinear method for the Gulf Stream region [Douglas et al., 1983].



Fig. 5. Dynamic ocean topography generated from a long-term mean sea surface consisting of 1.5 years of GEOS-3 and three months of Seasat altimeter data for the Gulf Stream region. A global geoid model (PGS-54) was subtracted from the altimetric surface to remove the contribution of gravity gradients [Cheney and Marsh, 1982].

estimate of the seasonal mean dynamic topography which can be related to that determined by satellite altimetry.

Geology, Bathymetry and Ice Mapping

Bathymetric features, including the mid-ocean ridges, trenches, fracture zones, plateaus, and seamounts all produce corresponding features in the marine geoid as observed by satellite altimetry. The larger features, for example the Pacific trench, are easily identified by inspection of the contour maps of the mean sea surface. Geoid features based upon hydrographic data, such as the GEM-9 geoid, are probably accurate to 10 cm at wavelengths greater than 10,000 km. Since this is comparable to the scale of most ocean basins, these geoid models might be used to determine the gyre-scale flows.

In the initial computation of the mean altimetric surface, the radial orbit error (which for Seasat and GEOS-3 is of the order of 1 m), must be treated. However, solutions have been generated which reproduce some features of dynamic topography maps based upon hydrographic data. Each one treats the radial orbit error in a different way. In the first case a global altimetric surface was computed from 1.5 years of GEOS-3 data combined with the 3-month Seasat data set. With such a large quantity of altimeter data over a relatively long period, much of the radial error was probably removed through averaging. When this surface is differenced with the PGS-54 geoid, a model developed especially for Seasat, gyre-scale features with the proper sense of flow are obtained (Figure 5). A global surface was also computed from only three

days of Seasat altimeter data from which a Fourier series representation of the radial orbit error was removed through an analysis of crossover differences (where ascending and descending ground tracks intersect), together with the along-track altimeter data. A third surface was determined from a spherical harmonic analysis of the differences between the GEM-9 geoid and a 3-month Seasat altimeter surface of the Pacific. All three of these surfaces show similar gyre-scale features, remarkable result considering the magnitude of the present altimetric and geoidal errors.

Jenkins briefly reviewed evidence that ³He is outgassed from the oceanic crust at a rate of about 6 × 10²⁰ atoms per year. A part of this flux is reflected in the high ³He concentrations near the East Pacific Rise at 20°S and 5°N, reflecting inputs from high temperature axial vents found along the crest of the EPR. The ³He distribution, along with the dynamical models of Stommel [1982], suggest an axial ³He flux from the EPR of about 0.6 × 10¹⁹ atoms per year. Assuming that the ³He/He ratio in all axial vents is 7.5 × 10⁻⁸ cal per atom [Jenkins et al., 1978], the heat flux for EPR axial vents is estimated by Jenkins as 0.5 × 10¹⁹ cal per year. A roughly equal axial flux might be expected from the remainder of the mid-ocean ridge system, giving a total high temperature axial heat flux of about 1 × 10¹⁹ cal per year. The remainder of the ³He is presumably degassed in convection systems on the ridge flanks.

Kathy Crane (Lamont-Doherty Geological Observatory) reported that she had towed thermistors over a 400-km segment of the Juan de Fuca Ridge, and a 300-km segment of the East Pacific Rise. Temperature anomalies suggested the presence of roughly one vent field for each 100-km of ridge length. Crane has used a dynamical model to calculate axial convective heat fluxes from the temperature data. A globally extrapolated value of 1.5 × 10¹⁹ cal per year is obtained, which is close to the upper limit calculated by Stommel [1982] assuming complete crustal quenching.

John Edmond (MIT) discussed hot spring chemistry in the three hydrothermal systems studied by his group, those at the Galapagos Spreading Center [Edmond et al., 1979], the East Pacific Rise at 21°N, and the sediment-covered ridge in the Guaymas Basin. In all systems, the discharging waters are depleted in Mg²⁺ and SO₄²⁻, and highly enriched in Li⁺, K⁺, and Rb⁺. Na⁺ and Cl⁻ concentrations range between approximately 90–110% of the seawater concentrations. H₂S concentrations are 4–9 mM. Values for pH and alkalinity are low or even negative for Galapagos and EPR, but (along with NH₄⁺) high in Guaymas owing to reactions with the sediment prior to discharge.

Assuming, for the sake of discussion, that the axial convective heat flux is equal to

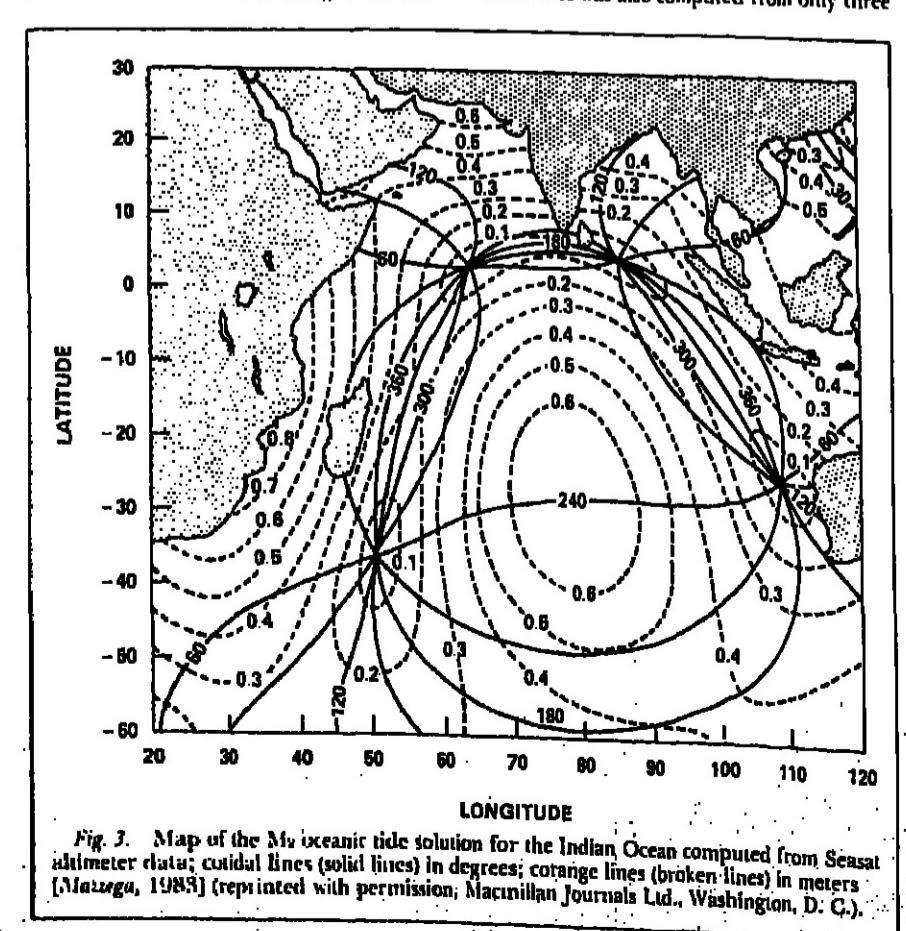


Fig. 7. Map of the M₂ oceanic tide solution for the Indian Ocean computed from Seasat altimeter data; solid lines (solid lines) in degrees; contour lines (broken lines) in meters [Mazzega, 1983] (reprinted with permission, Macmillan Journals Ltd., Washington, D. C.).

Article (cont.)

100 km have been resolved. In the Indo-Pacific basin, where sparse data coverage has limited previous tectonic studies, bathymetric features such as trenches, ridges, fracture zones and seamounts are clearly visible.

Seasat also tracked areas of smooth terrain including deserts, salt flats, ice sheets, tundra, and valleys. The altimeter tracker did not respond quickly enough over most non-ocean features; however, the waveform data have been retracted to achieve accuracy levels of better than a meter. The Seasat altimetric overland data base consists of more than 400 hours at a measurement rate of 0.1 second or potential overland profile lengths of approximately 107 km.

Overland analyses for south central Arizona, the Imperial Valley of California, the Yuma Valley of Arizona, and the Florida Everglades have yielded surface elevations over smooth terrain accurate plus or minus 1 m when correlated with large scale maps. Detailed analyses of ice topography in the polar regions have provided regional maps with a precision of a few decimeters.

Scientific Communications and Cooperation

The further scientific use of satellite altimetry may be fostered by joint geodesy/ocean sciences sessions at AGU meetings. There will remain the need for in-depth discussions of scientific results in a workshop format; the Pilot Ocean Data System Science Steering Group may sponsor the next altimetric workshop in 2 or 3 years.

European Missions

Since the Europeans will be active with their own altimetric satellites late in the decade, scientific communication between members of European and American scientific teams will be essential. Even more important is the promotion of freely-flowing, two-way data exchanges between these communities. **GEOSEAT**

Beginning in 1984, the Navy's GEOSAT promises to provide useful data for several years, bridging the GEOS-3/Seasat and the NRO/TOPEX eras, since it will be the only altimetric satellite available in that period, it is important that the ocean science community have access to the data at least in its unclassified ocean residual, i.e., after a classified geoid has been removed from the original data. Since there is likely to be extensive interest in this data set, it should be made broadly available through the Pilot Ocean Data System. However, there is a need to have adequate documentation and quality assurance of the classified aspects of the data stream in order to preserve the scientific integrity and utility of the data base.

Operational Applications

Synoptic sequences of altimetric topography on a weekly time scale could reveal the

ocean's space-time circulation structure as a function of mission parameters. Overall, the future use of satellite remote sensed data of the ocean, especially those associated with the ocean dynamical variables observed with the altimeter, will be critically linked to numerical models and scientific models and dynamical.

Algorithms should be developed that continuously output to the geophysical data records (GDR) the statistical uncertainties in the derived parameters. This is done with maximum-likelihood and error-propagation methods. These uncertainties then allow the user to decide whether and how to employ the GDR's.

An updated Seasat GDR should be produced. This GDR would have a more accurate orbit, improved geophysical and instrument corrections, and estimates of uncertainties.

Overland tracking by satellite altimeters represents an important new technique for terrain mapping, regional tectonic studies, monitoring of vertical crustal movements, and mapping ice sheet topography.

Scientific Communications and Cooperation

The seminar was organized by the Pilot Ocean Data System (J. G. Klose, Manager) and jointly sponsored by the Pilot Ocean Data System and the TOPEX Development Flight Project Office (C. A. Yanarone, Manager).

These activities are sponsored by the NASA Office of Space Science and Applications, Oceanic Processes Branch (W.S. Wilson, Chief) and the Information Systems Office (A. J. Villaseca, Manager).

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derstood at the present to constrain hydrothermal fluxes quantitatively.

The diversity of topics discussed in the symposium illustrates the wide range of information needed to understand and constrain hydrothermal fluxes. The component studies have generally been very successful in extending our understanding of the systems which have been studied. The great current challenge is to develop a methodology for using the results of detailed regional studies to constrain global heat, water, and chemical fluxes associated with deep sea hydrothermal processes.

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Oceanography (cont. from p. 83)

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This meeting report was contributed by Michael L. Bender, who is with the Graduate School of Oceanography, University of Rhode Island, Kingstown, RI 02881.

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Books (cont. from p. 85)

duces infiltration and causes runoff to be greater than that of an unleaded slope. This runoff enhancement caused by a decrease in permeability, known as "water harvesting" or "milking the hillside," obviously was understood by those who managed land and water in the ancient Negev. They could harvest perhaps not more than 25% of the seasonal rainfall. Modern technology can increase this percentage by sealing, waterprofiling, and stabilizing the soil cover. This source can be significant if one considers that 100 mm of rain on just 1 km² can produce 100,000 m³ (nearly 30 million gallons) of high-quality water.

This book is highly recommended to those readers who are interested in, or concerned with, deserts, people, or water. Fascination with the desert often borders on mysticism for those who know and love it. Perhaps our spiritual origins are in the desert in the same sense that our physical origins are in the sea. Time spent in the desert may represent a return to spiritual beginnings, rejuvenating the soul in much the same manner that time spent at the sea, the locus of our physical origins, rejuvenates the mind and body. One basic philosophy of the author is expressed in these few words on the last page: "Let us respect and love the desert, and seek to live with it, not rape or despoil it."

William Beck with the U.S. Geological Survey, Reston, VA 22092.



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Faculty Position/University of Massachusetts. Two tenure track faculty for new positions in September 1984 in a new doctoral program in Environmental Science. Successful candidates will provide evidence that they will conduct significant research programs, participate effectively in graduate education, and interact with other faculty in program development and administration. Rank will depend on experience and qualifications. Duties required: postdoctoral experience desirable. The appointments are to be at the Assistant Professor level, but applications from senior investigators will be considered. PHYSICAL OCEANOGRAPHY will develop graduate courses and a research program that will include studies of coastal, harbor, and estuarine dynamics and sedimentation processes. ENVIRONMENTAL POLICY AND LAW: To develop a core course and other courses, and a research program in legal/political aspects of environmental legislation, regulation and enforcement. These are extensive opportunities for collaboration with the John W. McCormack Institute of Public Policy. Applications should be submitted to the Department of Physics & Astronomy at the University of Iowa for qualified candidates with a Ph.D. degree and experience in space plasma and/or auroral physics. Present research in space plasma physics emphasizes analytical and numerical modeling of ionosphere-magnetosphere plasma interactions on board earth-orbiting spacecraft in the IMP and ISEE Missions. The University of Iowa's global imaging instrumentation on the spacecraft Dynamics Explorer 1 is the source of an extensive data base of auroral images from high altitudes at visible and ultraviolet wavelengths. Photometric observations are also available for other areas of research including the dynamics and sedimentation processes. The applicant should identify and describe areas of his or her expertise which can support experimental or theoretical investigations in space plasma physics and/or auroral physics. Salary and position will be determined by the applicant's qualifications and experience.

A resume and the names of three persons knowledgeable in the field of interest should be forwarded to: L. A. Frank, Department of Physics & Astronomy, University of Iowa, Van Allen Hall, Iowa City, Iowa 52242.

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Postdoctoral Research Fellow in Isotope Geology/University of Leeds, U.K. Applications are invited for a post of Postdoctoral Research Fellow, available immediately for a fixed term of 18 months.

Applicants should have experience in radiogenic isotope analysis and geochemistry, and interests in igneous geochemistry. The successful candidate will be expected to initiate research projects, collaborate on current research and assist in the training of research students in the isotope geology laboratories.

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A resume and the names

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The Department of Geology-Geography invites applications for a one year teaching position to replace a professor engaged in full time research. MS required; PhD preferred. Salary is commensurate with experience and credentials. Position starts September 1984. Teaching of introductory and upper division courses expected. Weber State College is a four year undergraduate institution located in Ogden, Utah. The Department has five geology faculty and six/tive majors. Application deadline: May 1, 1984. Send vita, transcripts and three letters of recommendation to: Dr. Wayne L. Wahlquist, c/o Personnel Department-1016, Weber State College, Ogden, Utah 84408.

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STUDENT OPPORTUNITIES

Opportunity for Graduate Study in Igneous Petrology/Isootope Geochemistry—Southern Methodist University. The Department of Geological Sciences at Southern Methodist University seeks outstanding individuals interested in a PhD program in igneous petrology and/or isotope geochemistry. The successful applicant should have a strong background in geology, chemistry, and mathematics and an interest in volcanic processes. Research will involve participation in a field-oriented petrological, geochemical, and isotopic study of Late

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Dr. M. A. Dunegan (214) 922-2782
Department of Geological Sciences
Southern Methodist University
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Research Assistantships/University of Maryland. The Meteorology Department of the University of Maryland has research assistantships available for graduate students. Fall Semester 1984. The Department offers courses of study leading to the degrees of Master of Science and Doctor of Philosophy in meteorology. Students with a Bachelor's degree in meteorology, the physical sciences, mathematics, or engineering are invited to apply.

Scholarships in the Maryland schools of Washington, D.C., the University is an ideal location for interaction with the large meteorological community of the area. The Department has cooperative research agreements with the National Oceanic and Atmospheric Administration and the National Aeronautics and Space Administration. Opportunities for internships and other government agencies, including the major centers at the National Center for Atmospheric Research and NASA, are important resources for students at Maryland. The Cooperative Institute for Climate Research and the Center for Ocean-Land-Atmosphere Interactions, both established recently on campus, offer the most expanded opportunities for research and study and research in climate analysis, modeling and prediction. A number of private and government agencies within the metropolitan Washington, D.C. area offer employment opportunities.

Interested individuals are encouraged to write to: Department of Meteorology, University of Maryland, College Park, MD 20742.

Research Fellowships at the University of Notre Dame. Fellowships in groundwater physics, hydrogeology, hydrogeophysics, and geochemistry are currently available in the Environmental Engineering Program of the Civil Engineering Department. Successful applicants will be awarded annual stipends of up to \$10,000/mo, plus full tuition. For additional information, contact Dr. Aaron A. Jennings, Department of Civil Engineering, University of Notre Dame, Notre Dame, IN 46556 (219) 239-4040.

State University of New York at Buffalo/Assistantship Opportunities. The Department of Geological Sciences invites graduate applicants for Fall 1984. Graduate assistantships are available at stipends up to \$5,500.00 for 10 months, plus tuition waiver. Special assistantships in geochemistry, geochemistry-mineralogy, and glaciology carrying a 10-month stipend of \$7,200.00 plus tuition waiver are available. Additional summer support is possible. Applications can be obtained from the Department of Geological Sciences, 1490 Ridge Lea, Amherst, NY 14228, (716) 645-3030. Deadlines for receipt of all materials is March 30, 1984.

The State University of New York at Buffalo is an affirmative action/equal opportunity employer and invites applications from minority and women candidates. No person in whatever relation with SUNYBuffalo shall be subject to discrimination on the basis of age, national origin, race, religion, or sex.

AGU Science and Policy on Capitol Hill

Arthur B. Weissman

I spent my year at the 1982-1983 AGU Congressional Science Fellow as legislative assistant to Sen. Christopher J. Dodd (D-Conn.), covering environment and energy issues. I offer this final report of my year not only for potential Congressional Science Fellows but for all readers of EOS who may want to contribute to science policy making (but are afraid to try).

My primary goal was to get a broad exposure to issues, organizations, people, and the legislative process; this I achieved beyond my expectation. Congress, remarkably, is set up like a vast marketplace: There are vendors (the lobby and interest groups) and consumers (the constituents), and each group makes its desires known to the members of Congress, who act as brokers. Issues are strewn about like so many fish—some stale and rancid; others fresh and meaty, the catch of the day. And there is no real respite. Problems requiring a response (if not a solution) will cut out even the most diligent Member of staffer.

Prospective applicants should have a broad background in science and be articulate, literate, flexible, and able to work well with people from diverse professional backgrounds. Prior experience in public policy is not necessary, although such experience and/or a demonstrable interest in applying science to the solution of public problems is desirable.

The fellowship carries with it a stipend of up to \$28,000, plus travel allowance.

Interested candidates should submit a letter of intent, a curriculum vitae, and three letters of recommendation to AGU. For further details, write Member Programs Division, American Geophysical Union, 2000 Florida Avenue, N.W., Washington, D.C. 20009 or telephone 462-6903 or 800-424-2488 outside the Washington, D.C. area.

Deadline: March 31, 1984

who think that science is irrelevant in policy making, there is the recent counterexample of budget director David Stockman holding up an acid deposition control program because it would interfere with the reauthorization of the Clean Air Act and the Clean Water Act and consideration of an acid deposition control program to name only a few. I watched in frustration as budgetary and defense matters marched through, and many of the issues on which I (and many others) worked hard and long remained stalled either in committee or on the floor. The lesson, though hard, was very clear.

Nonetheless, I gained considerable experience working on legislation and seeing what the legislative process entails. I developed some proposals for acid deposition control and some amendments on the ocean discharge waiver issue (Sec. 501(h) of the Clean Water Act), which directly affected Connecticut. In the course of drafting these bills for the Senator I learned how powerful can be many interest groups (helpful or troublesome, depending on their stand) and how powerfully Congressional committees guard their jurisdictions when noncommittee members offer proposals. I learned too how the scientific aspects of the issues, which occupied much of my time, played in counterpoint to other political considerations. And I saw how both science and politics got transmogrified by coexisting into cold, statutory language that elucidated neither the science nor the policies behind the measure.

There is much, however, that a scientist (and scientist) cannot do in Congress. Scientific information and reasoning in themselves do not provide usable answers to Congressional decision makers. Science must be transformed and cast in terms of social value judgments, which are the real fate of Congress. The acid deposition issue ultimately requires a balancing of regional and social costs and values. Environmental protection in one place may necessitate increased living costs or even job losses in another. What legislators need from scientists is a clarification of effect and consequences. For example, in the ocean discharge issue, scientists should clarify the possible effects on water quality of increased biological oxygen demand and toxic loading owing to reduced municipal sewage treatment.

Decisions on policy are value judgments based on many different considerations—social, economic, and political as well as scientific. The scientist may consider herself or himself equipped to render such judgments (still scientists are in fact asked to make such recommendations), but the scientist should realize that, in doing so, a professional boundary is being crossed. Charting the course of that fine line, and watching it bend and bulge with the flow of events, are among the imminent challenges and adventures of working as a scientist in Congress.

There is good news for the scientist who would influence legislation. With the plethora of issues to consider, much of the substance of issues and of the remedies proposed goes unexamined in Congress. Scientists can provide useful, substantive advice even if they do not work on the Hill, especially for the personal offices of Senators and Congressmen, which rarely have staff scientists at hand to help them with complex technical issues.

Concerned scientists should develop a relationship with the staff of their own Senator or Congressman, particularly with the legislative director and the legislative assistant in the relevant area. Correspondence is not sufficient (the Hill operates on personal contacts), so telephone calls and personal visits are always advised. An alternate route for influencing Congressional affairs is to become an advisor to an interest group. In either case, a persuasive argument communicated in a personal, direct way can be a significant influence on a Member of Congress (or his staff) looking for clarity in a maze of ideas, people, and paper.

Scientists (and science) can in fact do a lot of bad influence by scientists is the money Congress awarded last year to Catholic University and Columbia for materials science laboratories without any peer review or Congressional committee authorization. For those

Student Status

Theodore G. Apotria (T), D. Kent Backlin (T), William R. Blackport (H), William J. Dunn (GP), Steven Ecker (SS), Horacio F. Diaz (V), Brian Hauback (V), Robert W. Lane (H), Desmond Lee Ling Chye (S).

Stewart L. Moses (SM), Douglas Minell (G), Ee-Giang Pan (S), Daniel E. Sampson (V); Ingrid Sandhu (SM), Terry Seward (GP), James E. Stefano (T), Timothy Sukko (H), Sally Jo Stutts (T).

Gerald Tango (S), Joseph F. Tizer (S), Frank Toffoletti (SM), David M. Trull (S), Gail T. Vogt (T), Jonathan L. Wilson (S), Mark Zielinski (T).

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Meetings (cont. from p. 91)

ge and Geophysics, Schoepfstrasse 41, A-6020 Innsbruck, Austria) (Oct. 25-26, 1983).

June 1-4 International Seminar on Basin Strategy, Linkoping, Sweden. (U. Lohm, Water Theme, Linkoping Univ., S-58183, Linkoping, Sweden.) (Oct. 18, 1983.)

June 4-7 Seventh International Conference on Atmospheric Electricity, Albany, N.Y. Sponsors: IAPMA International Commission on Atmospheric Electricity, ANSI, and AGU. (R. E. Orville, Conference Chairman, U.S. Army Corps of Engineers, 1400 Washington Ave., SUNY, Albany, NY 12222; tel.: 518-437-3087.) (July 20, 1983.)

June 4-8 Third International Conference on Urban Storm Drainage, Göteborg, Sweden. Sponsors: IAHR and International Commission on Water Pollution Research. (P. Malmquist, Div. Dept. of Hydraulics, Royal Institute of Technology, 100-44 Göteborg, Sweden.) (July 20, 1983.)

June 6-8 Second American Conference on Ice-Nucleating Bacteria, Flagstaff, Ariz. (Ralph M. Billy Research Center, Box 6013, Northern Arizona Univ., Flagstaff, AZ 86011.) (Nov. 15, 1983.)

June 10-15 65th Annual Meeting of the American Association for the Advancement of Science (Physics Division), San Francisco, Calif. (John H. Vann, Dept. of Geography, California State Univ., Hayward, CA 94542; tel.: 415-881-3193.) (Jan. 31, 1984.)

June 11-12 Fifth European Conference on Environmental Pollution, Amsterdam, The Netherlands. (M. J. Blomberg, Dept. 1770, Cornwall, Ontario N1Y 1J7, Canada.)

June 11-15 Second Symposium on Critical Assessment of Forecasting in Western Water Resource Management, Seattle, Wash. Sponsors: AWRA and AGU. (G. R. Minton, President, Resources Planning Assoc., 113 Lynn St., Seattle, WA 98109; tel.: 206-282-1811.) (June 28, 1983.)

June 11-15 International Conference on Finite Elements in Water Resources, Burlington, Vt. Sponsors: Univ. of Vermont, AGU. (J. P. Laible, Dept. of Civil Engineering and Mechanical Engineering, Univ. of Vermont, Burlington, VT 05405; tel.: 802-656-3800.)

June 19-21 Third International Conference on Marine Simulation, Rotterdam, The Netherlands. Sponsored by Shell Nederland, with Institut Nauk i Techniki (Secretariat MARSH B), Maritime Research Institute Netherlands, P.O. Box 1555, 3000 RD Rotterdam, The Netherlands.)

June 23-30 Petrov Conference Melanges of the Appalachian Orogen, Newfoundland, Canada. Sponsors: GSA. (Brenna E. Lorenz, Dept. of Earth Sciences, Memorial Univ. of Newfoundland, St. John's, Newfoundland A1B 3X5 Canada.) (June 28, 1983.)

June 23-24 International Conference on Geomechanics, Denver, Colo. (A. L. Johnson, Woodward-Clyde Consultants, 7600 F. Orchard Rd., Englewood, CO 80111; tel.: 303-684-2770.)

June 24-26 International Symposium on Impenetrable Barriers for Soil Conservation, Colorado, A. L. Johnson, Woodward-Clyde Consultants, 7600 F. Orchard Rd., Englewood, CO 80111; tel.: 303-684-2770.)

June 23-30 14th International Conference on Mathematical Geophysics, Lecce, Italy. (L. Tronconi, INFINET/INGENIERIA, P.O. Box 31, Novara, Italy; tel.: 030-520000.)

June 25-27 Rock Mechanics in Protection and Production, 23rd US Symposium on Rock Mechanics, Evansville, Ind. Sponsors: AGU. (Charles H. Dowling, Dept. of Civil Engineering, Northwestern Univ., Evanston, IL 60201; tel.: 312-492-7070.) (Sept. 13, 1983.)

June 25-27 Seminar on Space Resources, 19th Congress on Space Resources, Graz, Austria. (Richard C. Hart, Space Science Board, 111-3028, National Academy of Sciences, 2001 Constitution Ave., N.W., Washington, DC 20230.)

June 25-July 7 Symposium on Space Observa-

tions for Climate Studies, Graz, Austria. Sponsor: World Climate Program. (S. Rutterberg, Secretary, COSPAR Commission A, NCAR, Boulder, CO 80307.) (Jan. 17, 1983.)

June 26-28 Symposium on the Achievements of the International Magmatic Study Group, Graz, Austria. Sponsors: ICSU Scientific Committee on Solar-Terrestrial Physics. (J. G. Rodger, Geological Institute, Univ. of Alaska, Fairbanks, AK 99701.)

June 26-28 International Conference on Atmospheric Electricity, Albany, N.Y. Sponsors: IAPMA International Commission on Atmospheric Electricity, ANSI, and AGU. (R. E. Orville, Conference Chairman, U.S. Army Corps of Engineers, 1400 Washington Ave., SUNY, Albany, NY 12222; tel.: 518-437-3087.) (July 20, 1983.)

June 28-30 International Conference on Urban Storm Drainage, Göteborg, Sweden. Sponsors: IAHR and International Commission on Water Pollution Research. (P. Malmquist, Div. Dept. of Hydraulics, Royal Institute of Technology, 100-44 Göteborg, Sweden.) (July 20, 1983.)

June 26-28 International Symposium on Deep Structure of the Mantle Crust, Results from Regional Seismology, Ithaca, NY 14853, Sponsor: AGU. (Mike Dungan, Dept. of Geological Sciences, Southern Methodist University, Dallas, TX 75275; tel.: 914-692-2760.) (Jan. 17, 1981.)

Aug. 26-29 Geothermal Resources Council 1984 Annual Meeting, Reno, Nev. (Geothermal Resources Council, P.O. Box 3500, Davis, CA 95721; tel.: 916-758-2360.) (Feb. 19, 1984.)

Aug. 26-31 Seventh Australasian Geological Convention, Sydney, Australia. Sponsors: Geological Convention Secretariat 7 AGC, P.O. Box 388, North Ryde, NSW, Australia 2113 (Nov. 29, 1983.)

Aug. 27-31 Seventh IAHR Symposium on Ice, Hamburg, Germany. (J. Schwartz, Ice Engineering Div., Hamburgische Schiffbau-Versuchsanstalt, FRG.) (Nov. 22, 1983.)

Aug. 27-Sept. 1 General Assembly of URSI, Florence, Italy. (Vito Capellini, IRE, 20171 Firenze, Italy; tel.: 051-501719.) (Dec. 1, 1983.)

Sept. 5-7 Quadrennial Ozone Symposium, Halkidiki, Greece. Sponsors: IAPMA International Ozone Commission (IOC), Commission of the European Communities, Dept. of Atmosphere, and WMO. (Christian Zerefos, Secretary, IOC, 10454 Athens, Greece.) (Sept. 5-7, 1983.)

Sept. 11-13 Annual Meeting of the American Geophysical Union, Washington, D.C. (John H. Vann, Dept. of Geography, California State Univ., Hayward, CA 94542; tel.: 415-881-3193.) (Jan. 31, 1984.)

June 11-12 Fifth European Conference on Environmental Pollution, Amsterdam, The Netherlands. (M. Blomberg, Dept. 1770, Cornwall, Ontario N1Y 1J7, Canada.)

June 11-15 Symposium on Critical Assessment of Forecasting in Western Water Resource Management, Seattle, Wash. Sponsors: AWRA and AGU. (G. R. Minton, President, Resources Planning Assoc., 113 Lynn St., Seattle, WA 98109; tel.: 206-282-1811.) (June 28, 1983.)

June 11-15 International Symposium on Space Techniques for Geodynamics, Szeged, Hungary. Sponsors: Hungarian Academy of Sciences and IAG/COSPAR Joint Commission on the International Coordination of Space Techniques for Geodesy and Geodynamics. (Ch. Reiter, Deutsches Geodätisches Forschungsinstitut, Abt. 1, Marburg, D-3550 Marburg; tel.: 06421-282, FRG.) (Nov. 15, 1983.)

June 19-21 13th International Symposium on Space Techniques for Geodynamics, Szeged, Hungary. Sponsors: Hungarian Academy of Sciences and IAG/COSPAR Joint Commission on the International Coordination of Space Techniques for Geodesy and Geodynamics. (Ch. Reiter, Deutsches Geodätisches Forschungsinstitut, Abt. 1, Marburg, D-3550 Marburg; tel.: 06421-282, FRG.) (Nov. 15, 1983.)

June 21-23 Seminar on Degradation, Removal, and Dispersion of Pollutants in Groundwater, Copenhagen, Denmark. Sponsor: International Association on Water Pollution Research and Control. (Eric Arvin, Dept. of Environmental Engineering, Technical Univ. of Denmark, Lyngby, Denmark.) (June 21-23, 1983.)

June 21-23 International Conference on Geodynamics, Denver, Colo. (A. L. Johnson, Woodward-Clyde Consultants, 7600 F. Orchard Rd., Englewood, CO 80111; tel.: 303-684-2770.)

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tion for the Study of Earth, Man, and the Moon, Dept. of Geological Sciences, Southern Methodist University, Dallas, TX 75275; tel.: 914-692-2760.) (Jan. 17, 1981.)

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Jan. 7-12 17th International Congress on Hydrogeology of Rocks of Low Permeability, Turin, Italy. Sponsors: International Association of Hydrogeologists, AGU. (F. S. Simpson, Dept. of Hydrology and Water Resources, College of Engineering, Univ. of Arizona, Tucson, AZ 85721; tel.: 602-621-6725; fax: 602-621-6726.) (Jan. 7-12, 1985.)

Aug. 27-Sept. 1 General Assembly of URSI, Florence, Italy. (Vito Capellini, IRE, 20171 Firenze, Italy; tel.: 051-501719.) (Dec. 1, 1983.)

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0745 Noise and Interference
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W. M. Godfrey (Formerly Department of Physics, University of California, Berkeley, presently Boeing MAC, Seattle, WA 98108) P. M. Houston, R. H. Koch, and J. Clarke
A reference seismometer array was used to measure the noise correlation lengths in remote reflections. The results show a significant difference between the noise correlation length and the noise correlation length in the seismic wavelet, and the noise correlation length in the seismic wavelet is about one-half the noise correlation length in the noise.

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